

1) Soil Forming Factors

- 2) Primary Properties that affect Water Movement in Soils
- 3) Soil Water Concepts
- 4) Where to get additional information

## **Today's Talking Points**

- Parent Material
- Climate
- Topography
- Biological Activity
- Time

# Soil Forming Factors

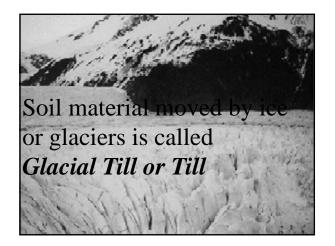
- Those that formed <u>in place</u> residuum
- Those that have been <u>transported</u> In Missouri: transported by Ice, Water, Wind and Gravity

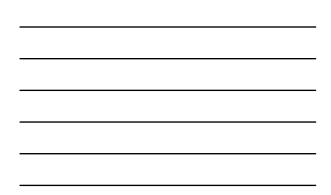
*Two Categories of Parent Material* 

Soil that form <u>in place</u> from the weathering of the underlying rock or minerals is called

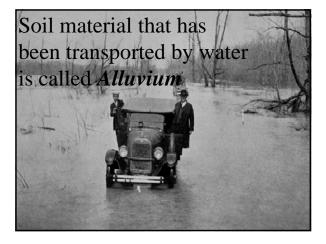
Residuum

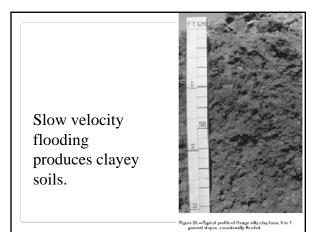




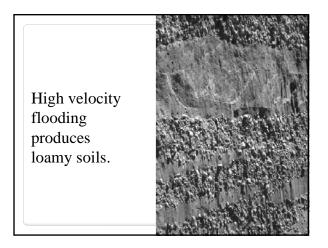












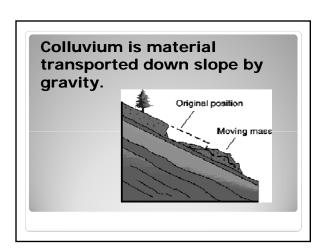


Wind deposited - the further from the source the loess gets thinner and has more clay (less silt)

Thickest deposits are adjacent to the Missouri and Mississippi River flood plains.

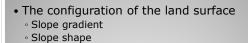
Loess



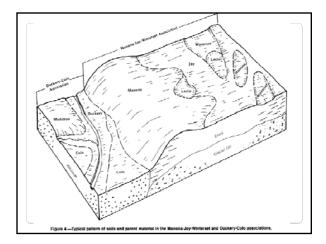


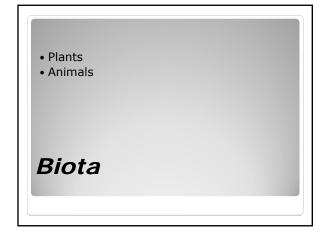
Temperature and Moisture are major influences. In the US, we have permafrost in Alaska and very warm conditions in our southern states. With regard to moisture, we have both deserts and rain forests.

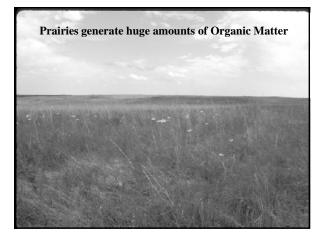
# Climate

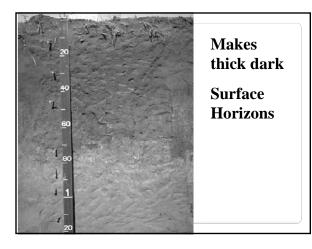


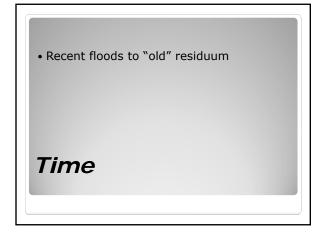
Topography









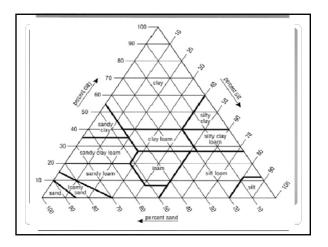




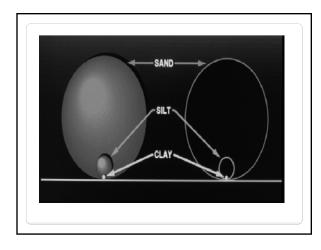
### 1) Soil Texture – particle size distribution

- 2) Soil Structure
- 3) Organic Matter Content4) Landform hill slope profile

#### **Primary Factors that influence** Water Movement through the Soil







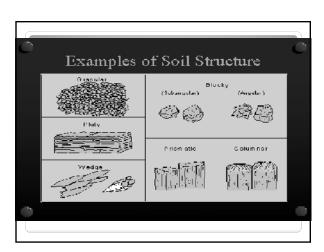


Particle Sizes Sand: 2mm to .05mm Silt: .05 mm to .002 mm Clay: smaller than .002mm

MLICA 4-9

- Blocky or granular surface layer, A horizons
- Platy or blocky subsurface, E horizons if present
  Blocky and/or Prismatic subsoil, B horizons (compound: prismatic parting to blocky)
- Massive or Single grained -substratum, C horizons

**Structure** 



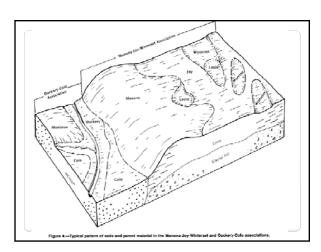
• Soil Organic Matter Content higher levels: • Will increase the soils ability to store and give up water plants.

• Will increase water movement rates through the soil.

#### **Soil Organic Matter**

- How we can describe shapes of the land surface.
- Slope shape and gradient are important considerations.

Landforms



In soil science, <u>permeability</u> is defined qualitatively as the <u>ease</u> with which <u>gases</u>, <u>liquids</u>, or <u>plant roots penetrate or pass though a soil mass</u> or layer. It is measured in length with no time component.
Saturated hydraulic conductivity is a quantitative measure of a saturated soil's ability to transmit water when subjected to a <u>hydraulic gradient</u>. It can be thought of as the ease with which pores of a saturated soil permit water movement. It is measured in both length and time.
Permeability classes and most references to permeability have been removed from the *NSSH* and replaced with the saturated hydraulic conductivity classes of the 1993 *Soil Survey Manual*.

Ksat Saturated Hydraulic Conductivity

SATURATED HYD	RAULIC CO	NDUCTI	VITY - PERMEA	BILITY	
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		141.14		20.00	
100.00	14.17		RAPID		
HGH		42.34		6.00	
nan			MODERATELY RAPID		
10.00	1.417	14.11		2.00	
MODERATE	LY	4.23	MODERATE	0.60	
HGH			MODERATELY SLOW	0.00	
1.00	0.1417	1.41		0.20	
MODERATE	LY	0.42	SLOW	0.06	
0.10	0.01417				
LOW			VERY SLOW		
0.01	naunnau	0.01		0.0015	
0.00 VERY LOW	V 0.00	0.00	IMPERMEABLE	0.00	
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µm/sec x 0.1417 = i	n/hr	in/hr x	7.0572 = µm/sec	_	



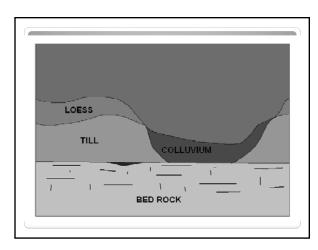
- Saturated hydraulic conductivity is the amount of water that would move vertically through a unit area of saturated soil in unit time under unit hydraulic gradient.
- If you find old permeability information, you can convert from *u*m/sec to in/hr by multiplying *u*m/sec by .1417

- Do you have an awareness of the factors of soil formation?
- How about the primary factors that influence water movement through soil?
- What is Saturated Hydraulic Conductivity?
- Ready to move on to Soil Water Concepts?

**Quick Review** 

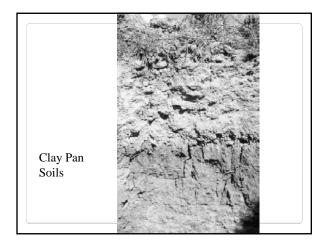
- Does water moves through all soils at the same rate?
- Soil horizons and restrictive features.
- Water tables.
- Interpretations: Hydraulic groups & Drainage classes.

### **Soil Water Concepts**



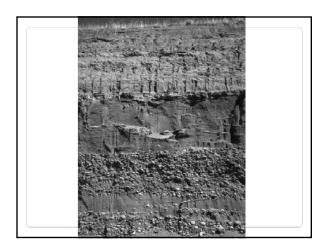
- clay pans
- abrupt texture change
- discontinuities
- bedrock

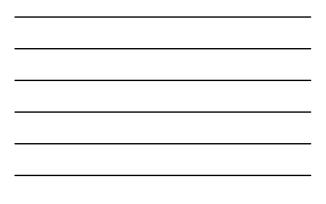
Soil Horizons and Restrictive Features







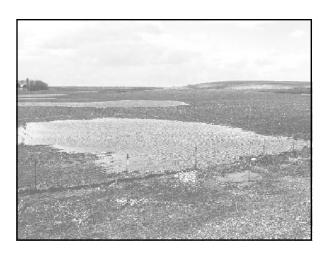


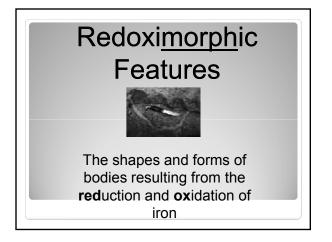


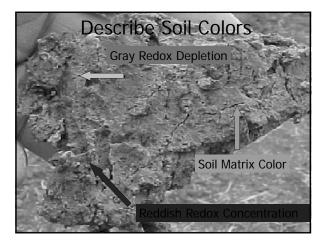
#### WATER TABLES

PERCHED (on something in the soil profile like) fragipans clay pans abrupt texture change discontinuities bedrock

APPARENT (flood plains) from the bottom up





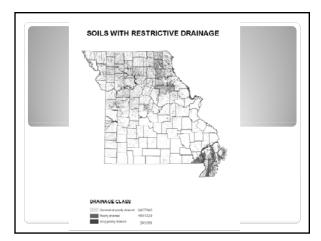




- Hydrologic Groups
- (a) Definition
- Hydrologic group is a group of soils having similar runoff potential under similar storm and cover conditions.

- (b) Classes
  The soils in the United States are placed into four groups, A, B, C, and D, and three dual classes, A/D. B/D, and C/D. In the definitions of the classes, inflictation rate is the rate at which water enters the soil at the surface and is controlled by the surface conditions. Transmission rate is the rate at which water moves in the soil and is controlled by soil properties. Definitions of the classes, and D, her at the surface and is controlled by the surface sollows:
  A. (Low runoff potential). The soils have a high infiltration rate even when thoroughly wetted. They chiefly consist of deep, well drained to excessively drained sands or gravels. They have a high rate of water transmission.
  B. The soils have a moderate infiltration rate when thoroughly wetted. They chiefly are moderately doep to deep, moderately well drained to well drained soils that have moderate are of water transmission.
  C. The solis have a slow infiltration rate when thoroughly wetted. They chiefly nave a layer that impedes downward movement of water transmission.
  D. (High runoff potential). The soils have a very slow infiltration rate when thoroughly wetted. They chiefly wave a layer that impedes downward movement of water transmission.
  D. (High runoff potential). The soils have a very slow infiltration rate when thoroughly wetted. They chiefly wetted they consist of clay soils that have a high swelling potential, soils that have a permanent of water transmission.
  D. (High runoff potential). The soils have a very slow infiltration rate when thoroughly wetted. They chiefly consist of clay soils that have a high swelling potential, soils that have a permanent high wetted they a soils that have a claypan or clay layer at or near the surface, and shallow soils over nearly impervious material. They have a very slow rate of water transmission.

• Dual hydrologic groups, A/D, B/D, and C/D, are given for certain wet soils that can be adequately drained. The first letter applies to the drained condition, the second to the undrained. Only soils that are rated D in their natural condition are assigned to dual classes. Soils may be assigned to dual groups if drainage is feasible and practical.





Definition.—"Drainage class" identifies the natural drainage condition of the soil. It refers to the frequency and duration of wet periods.

#### **DRAINAGE CLASS**

<ul> <li>Excessively Drained</li> <li>Somewhat Excessively Drained</li> <li>Well Drained</li> <li>Moderately Well Drained</li> <li>Somewhat Poorly Drained</li> <li>Poorly Drained</li> <li>Very Poorly Drained</li> </ul>	>6' >6' 3.5 to 6' 2-3.5' 1-2' 0-1' +1'
7 Drainage Classes use	d in MO

• Saturated hydraulic conductivity or Ksat Pertains to the amount of water that would move downward through a unit area of saturated in-place soil in unit time under unit hydraulic gradient.

 <u>Natural Drainage Class</u>
 Refers to the frequency and duration of wet periods similar to those under which the soil developed.

#### Ksat vs. DRAINAGE

- Air space size
- Air space quantity
- Changes impact physical processes in the soil. Such as adhesion, cohesion, adsorption, and surface tension.

Soil Air Space & Water Movement

- Compaction breaks down the natural soil structure.
- Compaction changes the air space size and quantity.

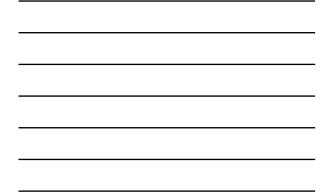
**Soil Compaction** 

 Saturated flow (also called gravitational flow) occurs only under <u>saturated</u> conditions when the force of gravity is greater than forces holding water in the soil.

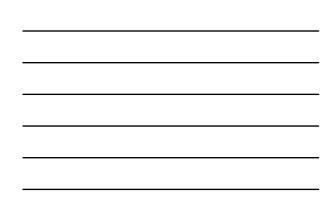
Capillary flow occurs in <u>unsaturated</u> soil (also called unsaturated flow). Unbroken films of water spread through connected capillary pores.

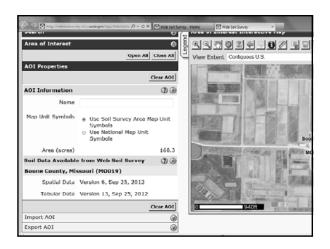
### Flow terminology



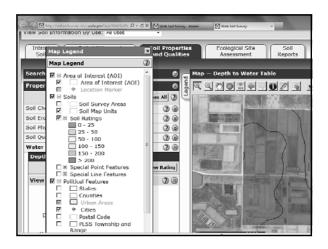


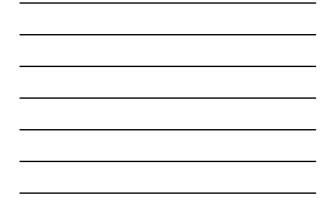
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50058–Mexico silt loam, 0 to 2 percent slopes								
Mexico	D	Very high	January	0.5-1.5	1.0-2.0	-	-	
	D	Very high	February	0.5-1.5	1.0-2.0	-	-	
	D	Very high	March	0.5-1.5	1.0-2.0	-	-	
	D	Very high	April	0.5-1.5	1.0-2.0	-	-	
	D	Very high	May	0.5-1.5	1.0-2.0	-	-	
	D	Very high	November	0.5-1.5	1.0-2.0	-	-	
	D	Very high	December	0.5 1.5	1.0 2.0			
50059—Mexico silt loam, 1 to 4 percent slopes, eroded								
Mexico	D	Very high	January	0.5-1.5	1.0-2.0	-	-	
	D	Very high	February	0.5-1.5	1.0-2.0	-	-	
	υ	Very high	March	0.5-1.5	1.0-2.0	-	-	



Boone County, Missouri							
Map symbol and soil name	Depth	Sand	Silt	Clay	Moist bulk density	Saturated hydraulic conductivity	Avai wa cap
	In	Pct	Pct	Pet	g/cc	micro m/sec	In/
50058—Mexico silt Ioam, 0 to 2 percent slopes							
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	15-34	1-2-5	30-39- 50	40-59- 60	1.25- 1.45	0.01-0.42	0.08
	34-42	1-2-5	55-66- 70	25-32- 38	1.40- 1.60	1.40-4.00	0.16
	42-80	4-4-10	40-66- 70	24-30- 15	1.40- 1.60	1.40-4.00	0.16
50059—Mexico silt loam, 1 to 4 percent slopes, eroded							
Mexico	0.7	5-9-12	65-71- 80	15-20- 26	1.40- 1.55	4.00-14.00	0.22
	7-12	5·9·12	55-61-	28-30-	1.45	1.40-4.00	0.21

#### SUMMARY:

- PRIMARY CONTROL OF WATER MOVEMENT IN SOIL: --Soil Texture – particle size distribution; Soil Structure; Organic Matter Content, Landform – hill slope profile.
- Saturated flow (also called gravitational flow) occurs only under saturated conditions when the force of gravity is greater than forces holding water in the soil.
- The more accurate and site specific, the more effective the drainage system.
- Soils are inherently complicated; therefore so is soil water movement.

MLICA 4-22

## Wetland Determinations

- Food Security Act USDA Program Participants.
- Clean Water Act Applies to all.

1. Name of Producer			2. I.D. Number (La	st 4 digits only) 3. Cu	renii Citoj	1 1 62
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<ol> <li>Are you a landlord on an If "TES", enter the farm</li> </ol>	y tam that will not be in compliance will number at the end of this statement	HELC and WC pro	vapors? County FSA Office before	completing this form:		
<ol> <li>Do say of your landlord a II "RES", under the larm</li> </ol>	elizes to comply with HELC requirement number al live end of this statement,	s on any farms? , or contact your C	ounly FSA Office before	completing this release		
7. List affiliated persons wit	h larning interests. See Page 2 for a	vexplanation. Ent	er "NONE", if applicable.			
					YES	N
<ol> <li>During the crop year enti- commodity on land for a</li> </ol>	ared in item 3 above, or the term of a re- hich a highly erodible determination has	puested USDA loan. I not been made?	did or will you plant or pro	duce an agricultural		
<ol> <li>On any land in which yo any activities (during the</li> </ol>	u have an interest, has anyone conduct current map year or the term of a map	sted any activities (s wind USDA loan) to	ince December 23, 1986)	or will anyone conduct		
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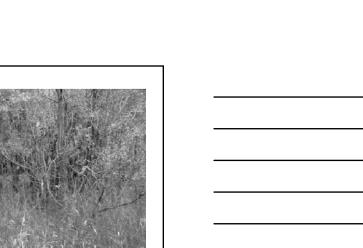
are:	8, or 10 }	"YES" for any one of these items, sign and date in item 11 below. A "YES" answer a this AD-1026 to NRCS to make a HELC and or certified wetland determinations. DO	NOT sign in Item 13
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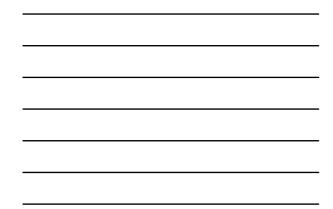
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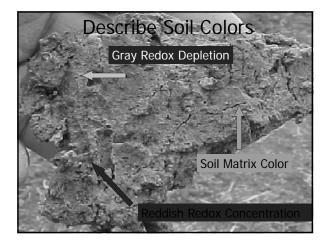


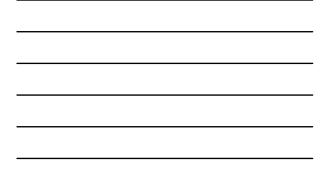
# Redoxi<u>morph</u>ic Features

The shapes and forms of bodies resulting from the **red**uction and **ox**idation of iron









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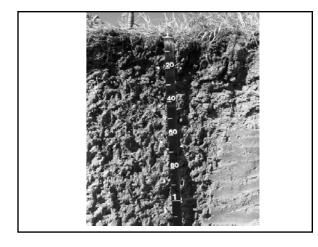
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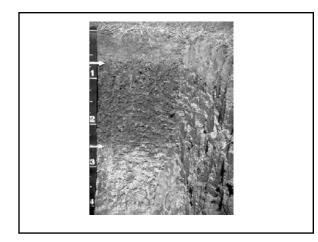




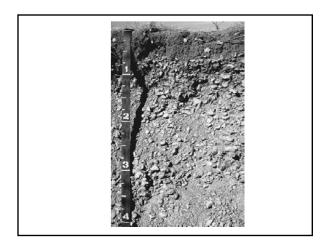




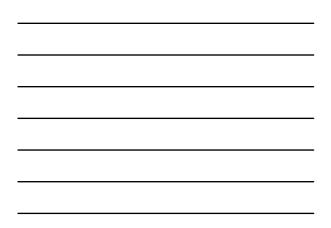












# Appendix 14D

# Auger-hole Procedure for Hydraulic Conductivity

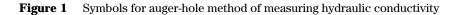
#### (1) Auger-hole method

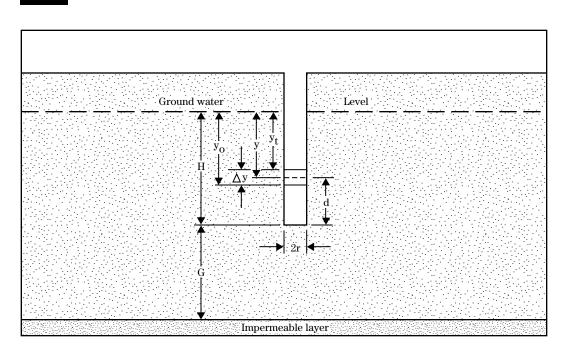
The auger-hole method is the simplest and most accurate way to determine soil permeability (fig. 1). The measurements obtained using this method are a combination of vertical and lateral conductivity, however, under most conditions, the measurements represent the lateral value. The most limiting obstacle for using this method is the need for a water table within that part of the soil profile to be evaluated. This limitation requires more intensive planning. Tests must be made when a water table is available during the wet season. Obtaining accurate readings using this method requires a thorough knowledge of the procedure.

The principle of the auger-hole method is simple. A hole is bored to a certain distance below the water table. This should be to a depth about 1 foot below the average depth of drains. The depth of water in the hole should be about 5 to 10 times the diameter of the hole. The water level is lowered by pumping or bailing, and the rate at which the ground water flows back into the hole is measured. The hydraulic conductivity can then be computed by a formula that relates the geometry of the hole to the rate at which the water flows into it. (i) Formulas for determination of hydraulic conductivity by auger-hole method—Determination of the hydraulic conductivity by the auger-hole method is affected by the location of the barrier or impermeable layer.

A barrier or impermeable layer is defined as a less permeable stratum, continuous over a major portion of the area and of such thickness as to provide a positive deterrent to the downward movement of ground water. The hydraulic conductivity of the barrier must be less than 10 percent of that of the overlying material if it is to be considered as a barrier. For the case where the impermeable layer coincides with the bottom of the hole, a formula for determining the hydraulic conductivity (K) has been developed by Van Bavel and Kirkham (1948).

$$\mathbf{K} = \left(\frac{2220\mathbf{r}}{\mathrm{SH}}\right) \left(\frac{\Delta \mathbf{y}}{\Delta \mathbf{t}}\right)$$
[1]





where:

- S = a function dependent on the geometry of the hole, the static depth of water, and the average depth of water during the test
- K = hydraulic conductivity (in/hr)
- H = depth of hole below the ground water table (in)
- r = radius of auger hole (in)
- y = distance between ground water level and the average level of water in the hole (in) for the time interval t (s)
- $\Delta y = rise of water (in) in auger hole during \Delta t$
- t = time interval (s)
- G = depth of the impermeable layer below the bottom of the hole (in). Impermeable layer is defined as a layer that has the permeability of no more than a tenth of the permeability of the layers above.
- d = average depth of water in auger hole during test (in)

A sample form for use in recording field observations and making the necessary computations is illustrated in figure 2. This includes a chart for determining the geometric function S for use in the formula for calculation of the hydraulic conductivity.

The more usual situation is where the bottom of the auger hole is some distance above the barrier. Formulas for computing the hydraulic conductivity in homogeneous soils by the auger-hole method have been developed for both cases (Ernst, 1950). These formulas (2 and 3) are converted to English units of measurement.

For the case where the impermeable layer is at the bottom of the auger-hole, G = 0:

$$\mathbf{K} = \frac{15,000r^2}{(H+10r)\left(2-\frac{\mathbf{y}}{H}\right)\mathbf{y}}\frac{\Delta \mathbf{y}}{\Delta t}$$
[2]

For the case where the impermeable layer is at a depth  $\geq 0.5$ H below the bottom of the auger hole:

$$K = \frac{16,667r^2}{(H+20r)\left(2-\frac{y}{H}\right)y}\frac{\Delta y}{\Delta t}$$
[3]

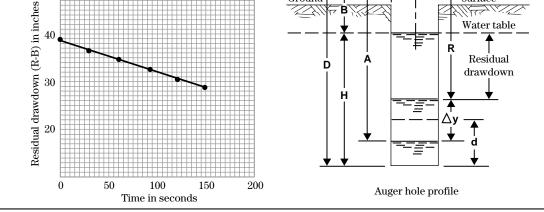
The following conditions should be met to obtain acceptable accuracy from use of the auger-hole method:

2r>2 1/2 and <5 1/2 inches H > 10 and < 80 inches y > 0.2 H G > H y < 1/4 y\_o

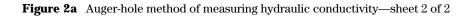
Charts have been prepared for solution of equation 3 for auger-holes of  $r = 1 \ 1/2$  and 2 inches. For the case where the impermeable layer is at the bottom of the auger hole, the hydraulic conductivity may be determined from these charts by multiplying the value obtained by a conversion factor f as indicated on figure 3.

Figure 2a Auger-hole method of measuring hydraulic conductivity—sheet 1 of 2

#### **Field Measurement of Hydraulic Conductivity** Auger-Hole Method For use only where bottom of hole coincides with barrier. Salt Flat Dry River Soil Conservation District Field Office Joe Doe - Farm No. 2 \_\_ Location\_\_1/2 Mi. E. Big Rock Jct. Cooperator\_ 264 4\_\_\_\_\_ Farm No.\_\_\_\_\_ B-817 SCD Agreement No. Field No. Tom Jones Date\_1 June 64 Technician Salinity (EC) Soil \_\_\_\_\_ Water\_ 5.6 1.0 in/hr 4 Boring No. \_\_ Estimated K\_\_\_ Distance to water surface Start from reference point Residual 10:03 ∆у During pumping Δt drawdown Before After Elapsed pumping pumping Time в Α R R-B A-R Seconds Inches Inches Inches Inches Inches Seconds ΧХ ΧХ 43 ΧХ ΧХ ΧХ ΧХ ΧХ ΧХ 81.5 0.0 0.00 38.5 30 79.0 36.0 60 77.5 34.5 90 33.0 76.0 120 74.0 31.0 150 150 72.0 9.5 29.0 50 – Ref. point Ground surface Water table



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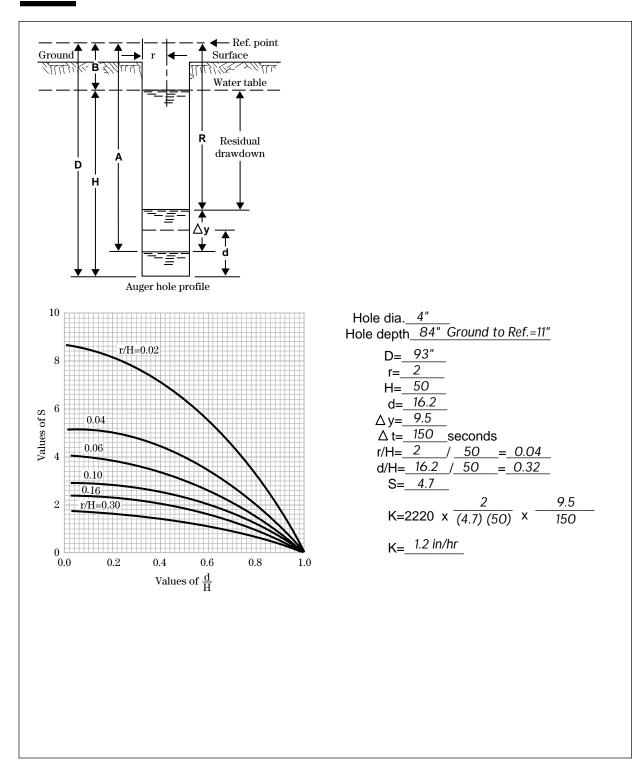
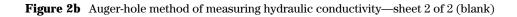


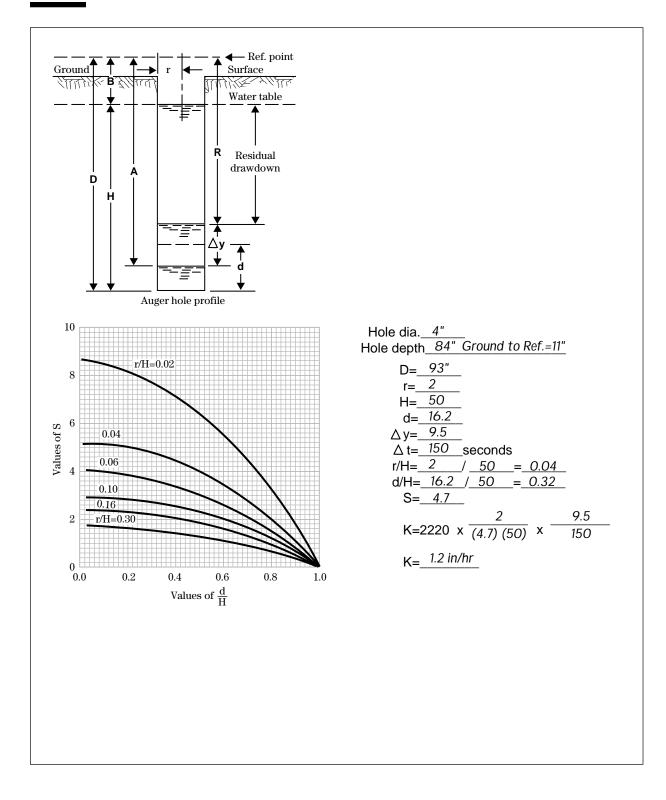
Figure 2b Auger-hole method of measuring hydraulic conductivity—sheet 1 of 2 (blank)

oil Conserva				e coincides with Field Offic			
ooperator							
CD Agreement NoFie			ld No Farm No				
echnician		Date					
oring No	Salini	ty (EC) Soil	Water_	Est	imated K		
Start	Distance to wat from reference					De status d	
10:03 Elapsed	∆t	Before pumping	After pumping	During pumping	– ∆y A-R	Residual drawdown R-B	
Time		В	A	R			
Seconds	Seconds	Inches	Inches	Inches	Inches	Inches	
50			Gro			Ref. poir surface	
in inches			—		-Ì <u>=_</u>	$-\frac{\text{Water table}}{\blacksquare}$	
R-B) i					-  ∓  - 	Residual	
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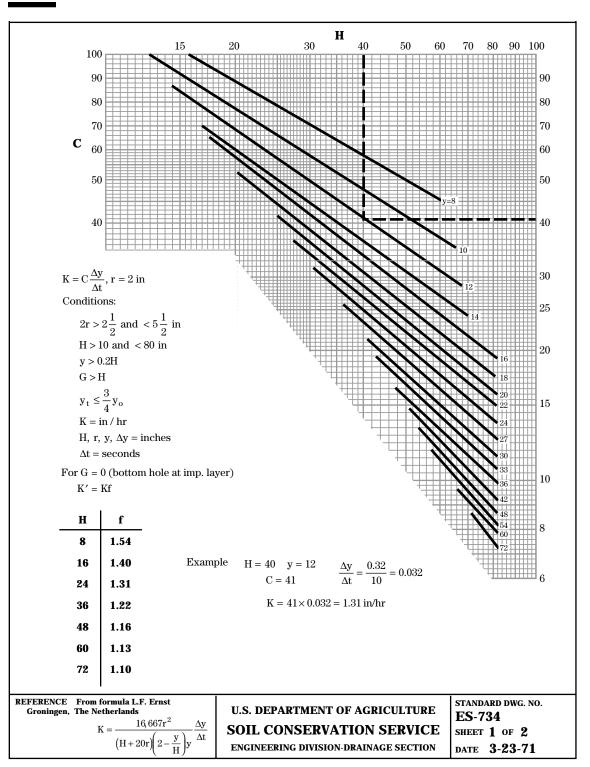
Water Management (Drainage)

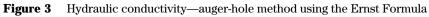
Part 650 Engineering Field Handbook





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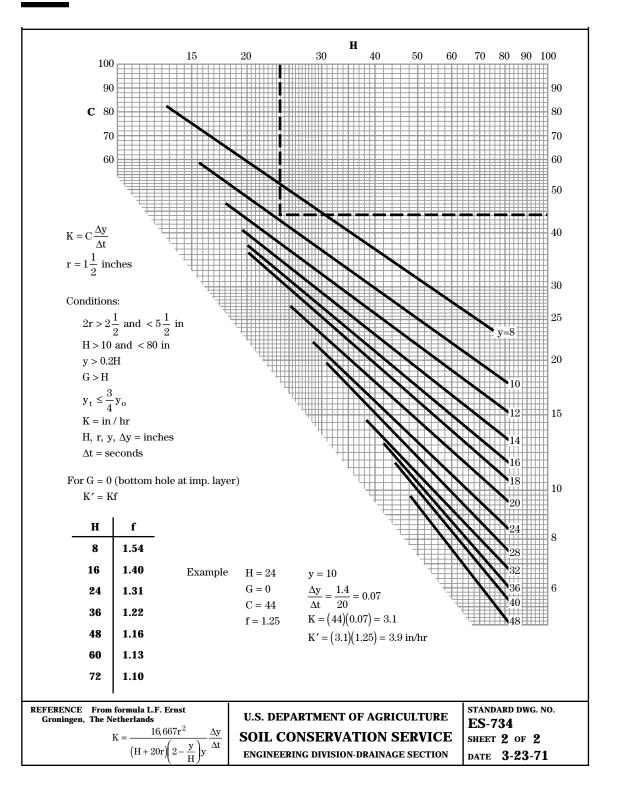




Water Management (Drainage)

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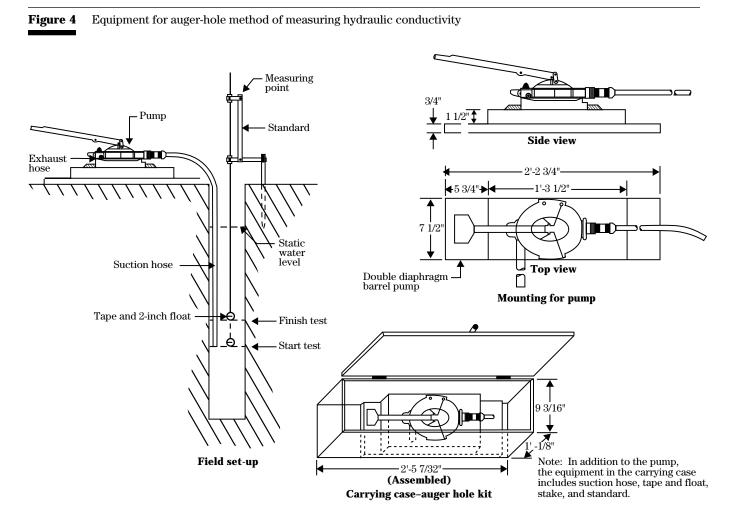
### *(ii) Equipment for auger-hole method*—The following equipment is required to test hydraulic conductivity:

- suitable auger
- pump or bail bucket to remove water from the hole
- watch with a second hand
- device for measuring the depth of water in the hole as it rises during recharge
- well screen may be necessary for use in unstable soils

Many operators prefer a well made, light weight boat or stirrup pump that is easily disassembled for cleaning. A small, double diaphragm barrel pump has given good service. It can be mounted on a wooden frame for ease of handling and use. For the depth measuring device, a light weight bamboo fishing rod marked in feet tenths and hundredths and that has a cork float works well. Other types of floats include a juice can with a standard soldered to one end to hold a light weight measuring rod.

A field kit for making the auger hole measurement of hydraulic conductivity is illustrated in figure 4. In addition to the items indicated in this figure, a watch and a soil auger are needed.

A perforated liner for the auger-hole is used in making the auger-hole measurement in fluid sands. This liner keeps the hole open and maintains the correct size. Several types of liners are used. Adequate slot openings or other perforations must be provided to allow free flow into the pipe.



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The openings in the screen should not restrict flow appreciably. The head loss through the screen should be negligible, and the velocity of flow through the openings should be small (0.3 foot per second or less) to prevent movement of fines into the hole. These criteria generally are met if the area of openings is 5 percent or more of the total screen area.

The Bureau of Reclamation uses 4-inch downspouting with 60 1/8- by 1-inch slots per foot of length. This works well in a variety of soils. A screen from the Netherlands is made from a punched brass sheet 2 millimeters thick with holes averaging about 0.5 millimeter in diameter. It is rolled into a tube 8 centimeters in diameter by 1 meter long. This screen works well because the sheet is rolled so that the direction in which the holes are punched is outward and the holes are variable in size. It has been used in many troublesome soils, and no clogging or failure to keep fines out of the hole has been reported.

Good judgment is needed in determining how far to drawdown the water level in the auger hole for the test. A minimum drawdown is necessary to physically satisfy theoretical criteria (refer to conditions given in fig. 3). Generally, a larger drawdown is made for slowly permeable soils than that for more permeable soils. A small drawdown for holes in sloughing soils may reduce the amount of sloughing. To prevent picking up sand in the pump, pumping should stop when the water level is within a few inches of the bottom of the hole.

Measurement of the rate of recovery of water in the auger hole should be completed before a fourth of the total amount of drawdown is recovered. Four or five readings should be taken at uniform short time intervals, and a plot of the readings made to determine a uniform rate of recovery to use in the formula. Plotting of time in seconds against the residual drawdown in inches indicates those readings at the beginning and end of the test that should be discarded and the proper values of t and y to use.

#### **Drainage Design Reference**

#### <u>Area</u>

1 acre = 43,560 square feet

#### **Saturated Hydraulic Conductivity**

1 micrometer per second = 1  $\mu$ m/sec

 $1 \mu m/sec = 0.2834$  feet per day

 $1 \,\mu\text{m/sec} = 0.1417$  inch per hour

1 inch per hour = 7.0572  $\mu$ m/sec

1 inch per hour = 2 feet per day

#### **Pipe Flow**

 $Q = V \times A$  and  $V = Q \div A$ 

Where: Q = Flow discharge rate, cubic feet per second

V = Flow velocity, feet per second

A = Cross Sectional Area, square feet

#### **Required Drainage Capacity**

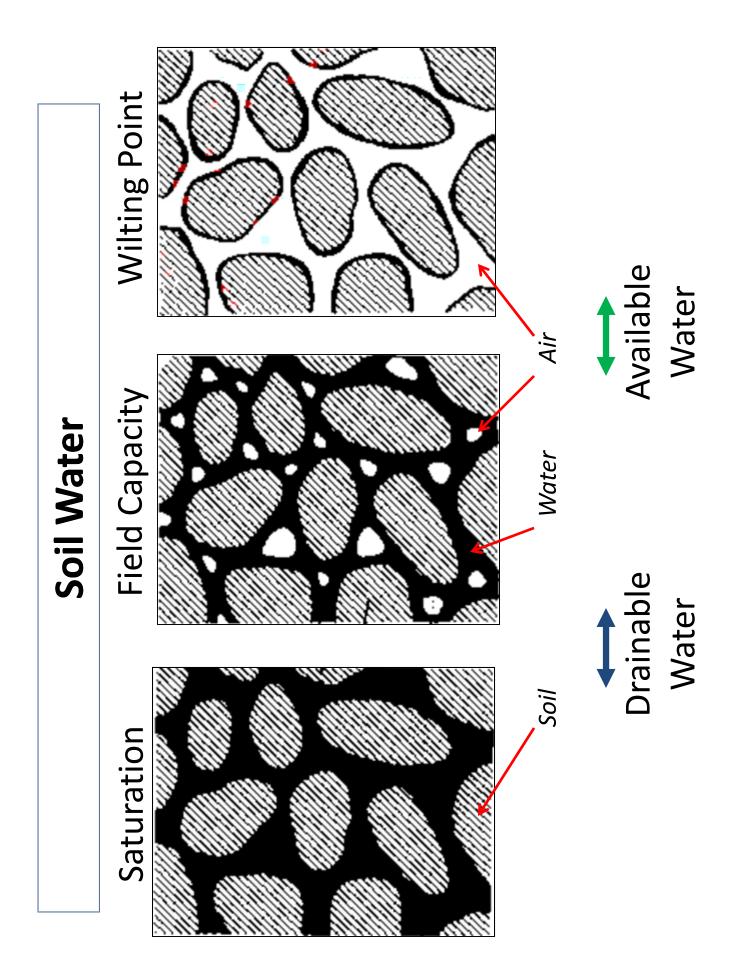
 $Q = 0.042 \times DC \times DA$ 

Where Q = Flow discharge rate, cubic feet per second

DC = Drainage Coefficient, inch/day

DA = Drained Area, acres

SATURATED HYDRAULIC CONDUCTIVITY - PERMEABILITY					
K <sub>sat</sub>	Class	Per	meability Cl	ass	
705.00	100.00	705.00		100.00	
VERY	HGH	141.14	VERY RAPID	20.00	
100.00	14.17		RARD		
		42.34		6.00	
HG		14.11	MODERATELY RAPID	2.00	
10.00	1.417		MODERATE		
MCCER/ HG		4.23	MODERATELY	0.60	
1.00	0.1417	1.41	SLOW	0.20	
MODERA		0.42	STOM	0.06	
0.10	0.01417				
Ь	N		VERY SLOW		
0.01	0.001417	0.01		0.0015	
VERY 0.00	LOW 0.00	0.00	IMPERIMEABLE	0.00	
µm/sec	in/hr	µm/sec		in/hr	
µm/sec x 0.1417			.0572=μm/sec		
$\mu m/\sec x 0.2834$	=ft/day	in/hrx2=ft/day			



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<ul><li>← Saturation</li><li>← Field Capacity</li></ul>			← Completely Dry	Soil Water Relationships (Typical for Clay, Clay Loam and Silty Clay)
	(Weak Capillary Forces) 13-21%	(Strong Adsorptive Forces) 15-24%		SOIL SOLIDS 45-65%
<b>€</b>	MUIOV e			IoV sbilos

**Typical Soil Water Relationships** 

Soil Texture	Wilting Point (% by vol.)	Available Water (% by vol.)	Drainable Water (% by vol.)
clays, clay loams, silty clays	15-24	15-26	3-11
well structured loams	8-17	12-22	10-15
sandy	3-10	7-20	18-35

Source: University of Minnesota BU-07644-S, Soil Water Concepts, Gary Sands

← Saturation			V Compretery Ury	Soil Water Relationships	
<b>DRAINABLE WATER 5%</b>	PLANT AVAILABLE WATER (Weak Capillary Forces) 20%	UNAVALABLE WATER (Strong Adsorptive Forces) 20%		SOIL SOLIDS 55%	
Solids Volume Pore Volume (Air and Water)					



with the goal of draining the top 12 inch layer in 48 hours. Given a soil (<u>silty clay</u>) with a drainable porosity of 5%

# Volume of drainable water

- = 5% x 12 inch depth
- = 0.6 inches

## Rate of removal = 0.6 inch ÷ 2 day = 0.3 inch/day



with the goal of draining the top 12 inch layer in 48 hours. Given a soil (<u>loam</u>) with a drainable porosity of 12%

# Volume of drainable water

- = 12% x 12 inch depth
- = 1.4 inches

## Rate of removal = 1.4 inch ÷ 2 day = 0.7 inch/day